

Rock sample descriptions as well as ana high-precision U-Pb zircon geochrono Bulletin B-45. Age results are given in Time Scale for relation of ages to periods	logy are presen millions of years	ted in the s ago (Ma).	accompanying
Rock unit and map symbol	<u>Age (Ma $\pm 2\sigma$)</u>	Sample	Map Locality
Pluton crystallization ages			
Agamenticus Complex, granite (Fb)	238.9 ± 0.3	12JAB4	G
Lyman pluton, biotite granite (Plg)	287.6 ± 0.4	MGS-032	В
Appledore diorite (within $O \in r$)	$361.1\pm\!\!0.4$	07SI-06B	L
Salamander Point diorite (Dd)	380.1 ± 0.5	MGS-038	Ι
Nisbitt Pond pluton, granite (Dnp)	$383.1\pm\!\!0.5$	MGS-035	А
Webhannet pluton, granodiorite (Dwgd)	382.9 ± 0.5	MGS-034	F
Webhannet pluton, granite (Dwg)	383.1 ± 0.5	MGS-033	Е
Biddeford pluton, granite (Dbig)	382.7 ± 0.5	ME-25	D
Biddeford pluton, granite (Dbig)	383.3 ± 0.5	ME-10	С
Breakfast Hill granite (Dbh)	402.9 ± 0.5	BH-1	К
Exeter pluton, diorite (Deg)	407.4 ± 0.5	89ex-7	Н
Age of youngest detrital zircon			

Cenozoic Era (Cz)	0-66	
Mesozoic Era (Mz)		
Cretaceous Period (K)	66-145	
Jurassic Period (J)	145-201	
Triassic Period (F)	201-252	
Paleozoic Era (Pz)		
Permian Period (P)	252-299	
Carboniferous Period (C)	299-359	
Devonian Period (D)	359-419	
Silurian Period (S)	419-444	
Ordovician Period (O)	444-485	
Cambrian Period $(\mathbf{\mathfrak{E}})$	485-541	
Precambrian time (p €)	Older than 541	
* In millions of years before	present. (Walker,	
J.D., Geissman, J.W., Boy	wring, S.A., and	
Babcock, L.E., compilers, 20	12 Geologic Time	
Scale v. 4.0: Geological So	ciety of America,	
doi: 10.1130/2012.CTS004R ²	3C.)	

BEDROCK GEOLOGY OF THE KITTERY 1:100,000 QUADRANGLE

the crystallization of masses of molten rock material (magma) that were entrapped within the Earth's crust. **STRATIFIED ROCKS**

Formation is isolated from the other formations by faults.

the east of the Merrimack Basin.

central New Hampshire and western Maine.

INTRUSIVE IGNEOUS ROCKS

following, listed approximately from oldest to youngest. possibly the Rochester plutons (**Dqd**). 2) Biddeford pluton (**Dbig**) is medium-grained biotite granite (**Photo 9**). Webhannet pluton consists of

occurring as isolated bodies or in migmatite in the Ridgemere Formation may be of the same age. Devonian age. 4) Lyman pluton (Plg) is fine-grained to medium-grained, irregularly textured biotite and biotite-muscovite granite. Irregular masses of granite pegmatite are common in many parts of the pluton. It is Permian, approximately 100 million years younger than most other granites in the Kittery 1:100:000 sheet.

also syenne so may be similar in age 6) Several relatively small gabbro complexes (Alfred, Cape Neddick, Tatnic, and Lebanon plutons) consi

shown by red symbols on the map. **DEFORMATION AND METAMORPHISM**

The stratified rocks have been extensively folded and faulted. Two stages of folding are particularly well Mine fault.

contact metamorphism of the adjacent rocks. **GEOLOGIC HISTORY** The oldest stratified rocks are those of the Rye Complex. Dating of tiny zircon grains in metasedimentary

perhaps multiple times. The Merrimack Group was deposited during Silurian time along the eastern edge of the Merrimack Basin, derived from peri-Gondwanan source areas to the east. Rapid deposition of sediments, a major regional deformation the source area had shifted to the east, probably the uplifted and deformed rocks of the Merrimack Basin. Acadian deformations included thrusting of the Merrimack Group over the Berwick and Shapleigh Group to the

zones are Alleghanian. The last major event recorded in the bedrock of the Kittery sheet is the intrusion of rocks of the White Mountain Plutonic-Volcanic Series of Mesozoic age, related to the opening of the Atlantic Ocean. These rocks comprise a northwest-trending chain of eroded volcanic centers from the New England Seamounts across central

Mesozoic age.

Blomshield, R.J., 1975, Superposed deformations on the Isles of Shoals, Maine-New Hampshire: M.S. thesis, University of New Hampshire, Durham, 54 p. Bothner, W.A., unpublished mapping. Society of America, Special Paper 331, p. 59-72. Hampshire, Durham, 127 p.

New England: Ph.D. dissertation, University of New Hampshire, Durham, 317 p. New Hampshire and Maine: M.S. thesis, University of New Hampshire, Durham, 128 p. enigmatic terrane: Atlantic Geology, v. 50, p. 138-154. University of New Hampshire, Durham, 218 p. M.S. thesis, University of New Hampshire, Durham, 114 p.

section: American Journal of Science, v. 287, no. 3, p. 242-264. Meeting, p. 15-28. Hussey, A.M., II, 1958-2004, field work. Hussey, A.M., II, 1961, Petrology and structure of three basic igneous complexes, southwestern Maine: Ph.D.

dissertation, University of Illinois, Urbana, 119 p. Hussey, A.M., II, 1962, The geology of southern York County, Maine: Maine Geological Survey, Bulletin 14, 67 p map scale 1:62,500. Hussey, A.M., II, 1980, The Rye Formation of Gerrish Island, Kittery, Maine: A reinterpretation: The Maine Geologist, Geological Society of Maine, v. 7, no. 2, p. 2-3. Hussey, A.M., II, Bothner, W.A., and Thompson, P.J., 2016, Bedrock geology of the Kittery 1:100,000 quadrangle, southwestern Maine and southeastern New Hampshire: Maine Geological Survey, Bulletin 45, 99 p. report, 83 figs., 3 tables. Lyons, J.B., Bothner, W.A., Moench, R.H., and Thompson, J.B., Jr., 1997, Bedrock geologic map of New Hampshire: U. S. Geological Survey, scale 1:250,000 (2 sheets). Novotny, R.F., 1969, The geology of the seacoast region, New Hampshire: New Hampshire Department of Resources and Economic Development, Concord, New Hampshire, edited by T.R. Meyers, 41 p., map scale 1:62.500. Novotny, R.F., unpublished notes and field maps related to his 1963 dissertation: Bedrock geology of the Dover-Exeter-Portsmouth region, New Hampshire: The Ohio State University, 182 p. Rickerich, S.F., 1983, Sedimentology, stratigraphy, and structure of the Kittery Formation in the Portsmouth, New Hampshire area: M.S. thesis, University of New Hampshire, Durham, 115 p. Schulz, J.E., 2004, Metamorphic petrology and structural implications of the calcareous (unnamed) member of the Berwick Formation within the Epping 7.5 minute quadrangle, southeast, New Hampshire: M.S. thesi University of New Hampshire, Durham, 108 p. Thompson, P.J., 2004, Bedrock geology of the Milton quadrangle, New Hampshire-Maine: Maine Geological Survey, Open-File Map 04-77, scale 1:24,000.

Bedrock exposed in the Kittery 1:100,000 quadrangle includes metamorphosed sedimentary rocks and intrusive igneous rocks. Sedimentary rocks are accumulations of particles of gravel, sand, silt, and clay in distinct beds, or precipitation of minerals from solution in layers or strata. Recrystallization of the sedimentary rocks by heat and pressure transforms them into metamorphic rocks, frequently developing minerals that are quite different from the originally deposited particles and changing the rock structure. Intrusive igneous rocks are those rocks formed by

Stratified, or layered rocks in the Kittery quadrangle are metamorphosed deposits of shale, siltstone, and sandstone that accumulated in two ancient ocean basins, the Merrimack Basin to the east and the Central Maine Basin to the west. (Note that present day orientations are used here for reference, even though the continents have moved significantly since the basins formed in Paleozoic time.) Deposition occurred during latest Ordovician through Middle Silurian time in the Merrimack Basin, and from latest Ordovician through Early Devonian time in the Central Maine Basin. These rocks were derived from weathering and erosion from exposed landmasses surrounding the basins. Stratified rocks are separated into formations of closely similar rock types that can be mapped over a broad area. Within the Kittery quadrangle these formations include the Eliot (Se) and Kittery (Sk) formations deposited in the Merrimack Basin, and formations of the Shapleigh Group (Rindgemere, Sr, Gully Oven, Sgo, Towow, St, and East Rochester, Der) deposited in the Central Maine Basin. The Berwick Formation (DSb) and the Phyllonite at Church Road (Scr) may have been deposited also in the Central Maine Basin, but the Phyllonite at Church road is so strongly deformed that its original rock is difficult to recognize, and the Berwick The Eliot Formation consists of metamorphosed thin beds of alternating tan-weathering limy siltstone, and dark gray shale (Photos 1 and 2). The Kittery Formation consists of variably thin to very thick-bedded metamorphosed limy sandstone, siltstone, and shale that preserve a variety of sedimentary features similar to those found today in deep-sea sediments that accumulate at the base of continental slopes. Graded bedding (Photo 3) and flute casts (**Photo 4**) are common and enable us to determine the stratigraphic order and the direction from which the sediments were derived and transported. The Kittery Formation was derived from a continental land mass lying to The Berwick Formation consists of metamorphosed calcareous sandstone, siltstone, and shale much like the Kittery Formation but with generally thicker beds (Photo 5). The Rindgemere, Gully Oven, and East Rochester Formations consist of gray, metamorphosed shale, siltstone, and sandstone. Bedding of the Rindgemere Formation is weakly developed and variable, and primary sedimentary structures are rare. On the contrary, bedding in the Gully Oven (Photo 6) and East Rochester (Photo 7) formations is characteristically well developed, with rhythmic alternations of shale and siltstone which are commonly graded. The Gully Oven Formation includes thin beds of reddish garnet-quartz granofels that help identify the formation (Photo 8). The Towow Formation is a distinctive alternation of rusty-weathering, sulfidic, poorly bedded to massive metamorphosed shale and siltstone. The Merchants Row Grits (Stmr) occurs locally at the top of the Towow Formation, characterized by granule-sized clasts up to 4 mm in size consisting of white and blue quartz, quartzite and rare metashale chips, perhaps representing a minor episode of erosion. Between the Towow and the East Rochester Formations lies an unnamed thin, possibly discontinuous, unit (DScb) of metamorphosed calcareous sandstone, similar to a widespread unit in

The intrusive igneous rocks are generally even-textured, massive rocks with interlocking mineral grains. They are distinguished from each other by characteristics such as the size of the mineral grains (texture) and proportions of minerals (mineralogy). Intrusive rocks occur in bodies (called plutons) that range in size from large areas shown on the map to small bodies intruding other rocks within an outcrop. The major bodies include the 1) Exeter pluton (Deg) is medium-grained variably textured quartz-biotite diorite with minor irregular masses of gabbro to quartz monzonite. Similar small bodies of diorite (Dd) are also considered Early Devonian, and

three principal phases. The oldest is fine-grained to medium-grained biotite-hornblende granodiorite (Dwgd); the intermediate-aged phase is medium-grained to coarse-grained biotite granite with moderate amounts of sphene and epidote (Dwg); the youngest phase, forming the northern end of the pluton, is medium-grained biotite-muscovite granite (Dwb). The Nisbitt Pond pluton (Dnp) is biotite-muscovite granite with white feldspar grains up to 3 cm long, commonly aligned, which form a swirled foliation pattern in many outcrops. Ages obtained for all of these bodies are the same, at the Middle-Late Devonian boundary. Many bodies of similar granite (Dg) and pegmatite 3) Barrington pluton (Dbag) is medium-grained to coarse-grained biotite-muscovite granite of Late

5) Agamenticus Complex is a roughly circular body consisting of several phases of pink, gray, or green granite and syenite closely associated in space and time. Syenite is similar to granite but with only small amounts of quartz. Several phases of granite and syenite contain very unusual varieties of dark minerals including riebeckite, arfvedsonite, enigmatite, aegirine-augite, and ferrohastingsite in place of biotite or hornblende of more typical granite. This reflects a higher than usual proportion of sodium in the magmas from which the rocks formed, a composition referred to as alkalic. The youngest phase of the complex is pinkish gray biotite granite of Triassic age (**Fb**) that cuts across the other phases of the complex. The Chase Stock (Mzc), known only from sea floor grab

primarily of gabbro and other dark gray to black, iron- and magnesium-rich igneous rocks. These are the youngest crustal rocks in the Kittery quadrangle, having intruded about 120 million years ago (Ma) in the Cretaceous Period. 7) Thin dikes of basalt intrude all the metamorphosed stratified rocks and all the igneous plutons except the Biddeford Pool (Photo 16). A few larger occurrences are mapped (KJd), and some representative examples are centimeters thick.

displayed in the Kittery Formation along the coast between Ogunquit and Kittery. Earlier folds are recumbent, that is, they are lying over on their side (Photo 10), whereas the later folds are upright, with both sides of the fold dipping steeply (**Photo 11**). Major faults in the map sheet include: the Phyllonite at Church Road (**Photo 12**), a shear zone which defines the boundary between the Shapleigh Group and the Berwick Formation; the Calef-Nonesuch River fault, separating the Berwick from the Merrimack Group; several faults cutting the Merrimack Group near the Great Bay, New Hampshire; the Portsmouth fault separating the Merrimack Group and Rye Complex; and the Great Common fault within the Rye Complex. Old silver mines and prospect pits of the Acton Silver District are located along a zone of white bull quartz, apparently marking the trace of a northwest-trending Mesozoic fault, the Silver Most of the large scale folding is probably the result of the Acadian orogeny, a major period of middle Paleozoic crustal compression driven by plate tectonic movements. An additional effect of this event was the recrystallization of the stratified rocks to form schists out of shales and siltstones, and granofels and gneisses from sandstones. The degree of heating during this compression is expressed in the variety of new minerals that formed, particularly in the schists. Lower temperature, for example, is indicated by the presence of chlorite, medium although the regional metamorphic pattern is complicated. Heat carried by intrusion of the large plutons caused formation.

rocks of the Rye indicates an origin as sedimentary and perhaps minor volcanic(?) rocks (Photos 13 and 14) that accumulated in Cambrian-Early Ordovician time. We suspect this assemblage was deposited on Ganderian basement. The Rye Complex was metamorphosed and deformed (Photo 15) during and after Ordovician intrusion,

(Salinic), and associated metamorphism are strongly implied before post-tectonic intrusion of the Exeter pluton in Early Devonian time. To the west, in the Central Maine Basin, sediments of the Shapleigh Group accumulated in Silurian to Early Devonian time. Silurian and older formations were derived from the Laurentian continent and Bronson Hill island-arc complex to the west. By the time of deposition of the East Rochester Formation, however, Rocks of the Central Maine Basin were multiply deformed and metamorphosed during the main Acadian event and intruded by Devonian plutons such as the Biddeford, Webhannet, Nisbitt Pond, and Barrington. The

pluton in the north part of the Kittery 1:100,000 map sheet, probably related to the Sebago batholith to the north of Permian age. This age suggests a relationship to the Alleghanian orogeny, the terminal compressional event of the Appalachian Mountains that formed the supercontinent Pangea. Although major deformational features of this age have not been confirmed in the area of the Kittery map sheet, it is possible that some of the major faults and shear

New England into southern Quebec. The rift and drift-related magmatism spanned more than 100 million years. The Agamenticus ring complex is of Triassic age, one of the oldest intrusive bodies of the series. The last vestiges of this activity are represented by the Cretaceous gabbroic complexes. In addition to these complexes, the area is crosscut by abundant basalt and diabase dikes (Photo 16), some of them tens of meters thick and traceable for tens of kilometers across the map area, and hundreds of kilometers to the northeast and southwest. They maintain an orientation roughly parallel to the edge of the North American plate where it split and separated from the Eurasian plate. A majority of these dikes are of Jurassic age, intruded after the Agamenticus Complex but before the gabbro complexes. Minor faults, such as the Silver Mine fault and some of the faults in New Hampshire may also be of

REFERENCES

Bothner, W.A., and Hussey, A.M., II, 1999, Norumbega connections: Casco Bay, Maine, to Massachusetts? In Ludman, Allan, and West, D.P., Jr. (editors) Norumbega fault system of the northern Appalachians: Geological Brooks, J.A., 1986, Bedrock geology of New Hampshire's inner continental shelf: M.S. thesis, University of New Brooks, J.A., 1990, The petrogenesis of the Agamenticus Complex and late Paleozoic and Mesozoic tectonics in shoreline. Carrigan, J.A., 1984, Geology of the Rye Formation: New Castle Island and adjacent areas of Portsmouth Harbor, Dorais, M.J., Bothner, W.A., and Buchwaldt, Robert, 2014, The Appledore Island pluton of the Rye Complex, coastal New Hampshire and Maine, USA: geochronological and chemical evidence for the affinity of an Escamilla-Casas, J.C., 2003, Bedrock geology of the seacoast region of New Hampshire, U.S.A.: Ph.D. dissertation, Eusden, J.D., 1984, The bedrock geology of part of the Alton, New Hampshire and Berwick, Maine 15' quadrangles: Eusden, J.D., Jr., Bothner, W.A., and Hussey, A.M., II, 1987, The Kearsarge-central Maine synclinorium of southeastern New Hampshire and southwestern Maine; stratigraphic and structural relations of an inverted Fargo, T.G., and Bothner, W.A., 1995, Polydeformation in the Merrimack Group, southeastern New Hampshire and southwestern Maine. In Hussey, A.M., II, and Johnston, R.A. (editors), Guidebook to field trips in southern Maine and adjacent New Hampshire: New England Intercollegiate Geological Conference, 87th Annual



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Home page: http://www.maine.gov/dacf/mgs/

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Geologic Map 16-6 2016 (Corrected 2019) For additional information, see Bulletin 45 (Hussey and others, 2016). This map supersedes Geologic Map 08-78.

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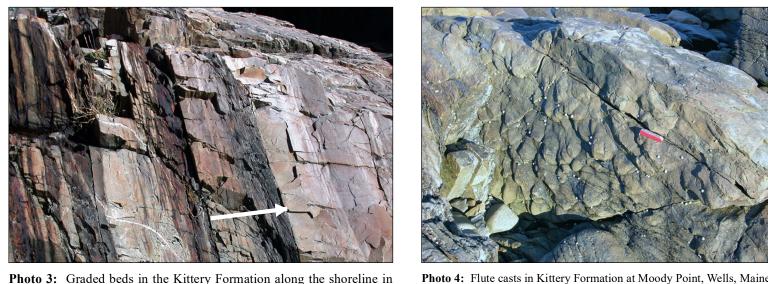
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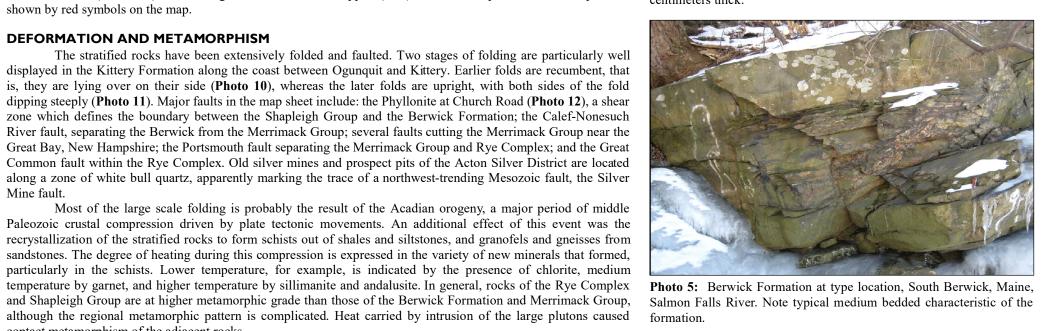
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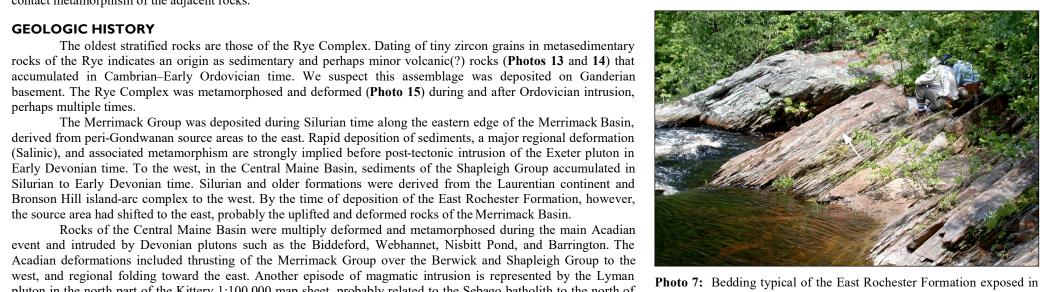
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York, Maine. A graded bed (two of which are well shown) has a light gabbro complexes and the biotite granite of the Agamenticus Complex (**Fb**). They occur in thin, nearly vertical gray metasandstone lower part that grades upward into a dark sheets ranging from a few centimeters to 75 meters in thickness. The dikes are black to very dark green, gray inclusion for a few centimeters to 75 meters in thickness. The dikes are black to very dark green, predominantly basalt and diabase (a coarser grained variety of basalt), approximately the same composition as These beds top to the right of the photo and are inclined because of gabbro. They are especially common in the exposures of the Kittery Formation along the coast from Kittery to folding. The graded bed marked by the arrow is approximately eight





indicates these beds are overturned.

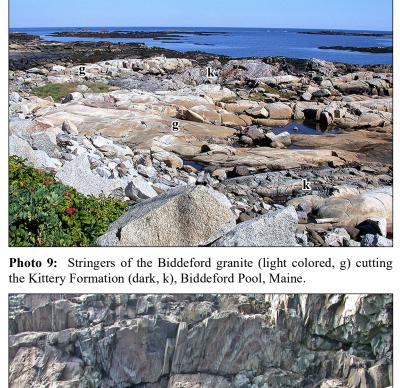


Photo 11: Upright folds of the Kittery Formation along York, Maine,

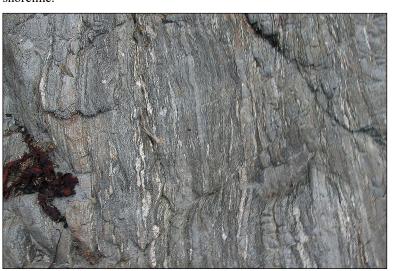
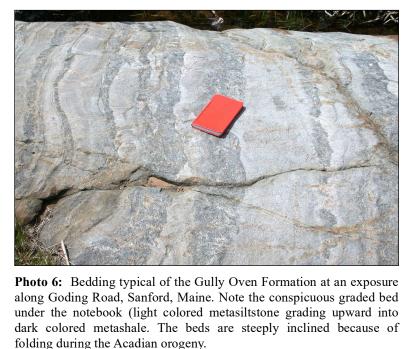


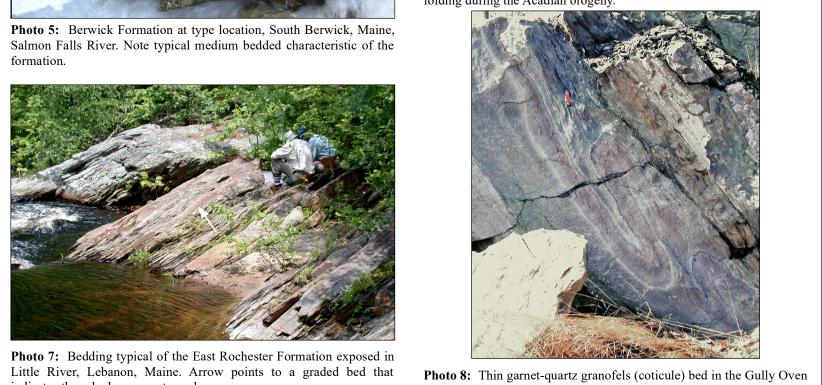
Photo 13: Sheared beds of metasiltstone (medium gray) and metashale Photo 14: Thinly laminated amphibolite, possibly representing mafic (dark gray) typical of the Rye Complex at Fort Foster, Kittery, Maine. volcanic rocks in the Rye Complex at Sewards Point, Kittery, Maine.

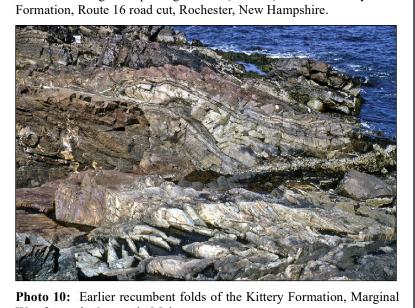


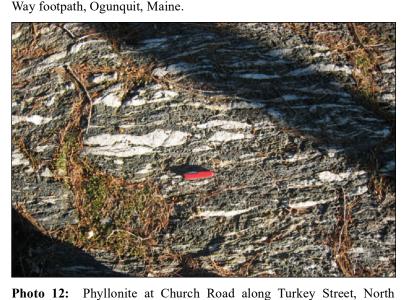
Hampshire, Coast Guard station. The blastomylonitic texture in the gneiss is the product of extreme early shearing. This exposure also contains late pseudotachylite (frictional melt rock) in S-shaped gashes formed by late shearing.





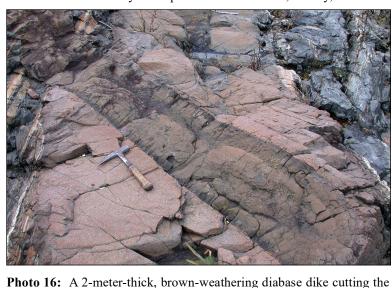








Berwick, Maine.



Kittery Formation, in turn cut by a thinner basalt dike. Shoreline north